

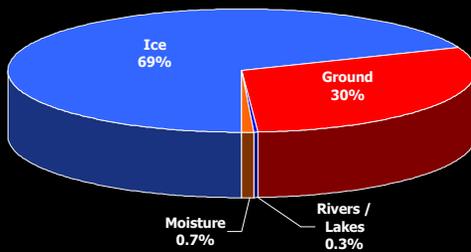
# Groundwater Hydrology

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## Why study groundwater?

- Important source of fresh water
- Large portion of the Earth's fresh water as groundwater
- Important to know the renewal time
  - Residence time  $\sim 5000$  a
- Self-sustaining use important (slowly renewable source)

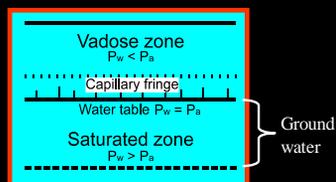
## Fresh water on Earth



## What is groundwater?

- If a hole is dug the water that flows freely into the hole is groundwater.
  - Since the air in the hole is at atmospheric pressure, the pressure in the groundwater must be higher if it is to flow freely into the hole.
- The zone between ground surface and the top of the groundwater is called the *vadose zone* (unsaturated zone).
  - Water there is held to the soil particles by capillary forces.

## Water beneath land surface



## Aquifers / Veitar

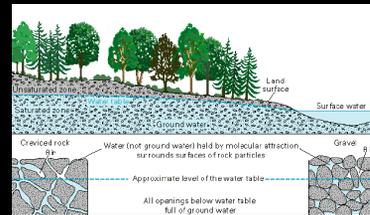
- Groundwater-bearing formations sufficiently permeable to transmit and yield water in usable quantities are called aquifers.
- Two types of aquifers:
  - (1) unconfined and (2) confined.

## Source of Groundwater

- **Atmospheric precipitation** is the main source of fresh groundwater.
- The water may have *infiltrated* directly into the ground where it landed, or been *collected* via *surface runoff* and then seeped into the ground.

## Groundwater system

Grunnvatnskerfi



## Vatn streymir úr hlíðum

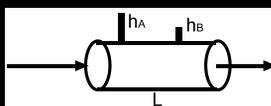


## Vatnafræði GROUNDWATER FLOW

Darcy

## Pipe flow – Darcy's law

- Demonstration of Darcy's law
- Pipe filled with sand
- Water applied under pressure at A
- Discharges at B



## Darcy's law

$$q = \frac{Q}{A} = -K \frac{dh}{dl}$$

- $q$  – specific discharge, [ $L T^{-1}$ ]
- $Q$  – discharge, [ $L^3 T^{-1}$ ]
- $A$  – cross sectional area, [ $L^2$ ]
- $dh/dl$  – hydraulic gradient
- $K$  – hydraulic conductivity [ $L T^{-1}$ ]

### Hydraulic conductivity

- Also referred to as the **Coefficient of permeability**
- Intrinsic permeability,  $K_i = C d^2$ 
  - C – shape factor
  - d – grain size (diameter)
- Hydraulic conductivity then,  $K = K_i * \gamma / \mu$ , where  $\gamma = \rho g$  and  $\mu$  is dynamic viscosity.

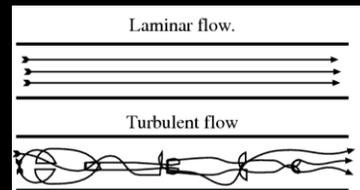
### Velocity of the water

- v sometimes called Darcy velocity  $v = \frac{Q}{A}$
- Note that is the average velocity through the cross section, not the water velocity
- Average velocity of the water itself  $v_x = \frac{Q}{n_e A}$

### Applicability of Darcy's law

- Only for laminar flow, not turbulent
- Reynolds number used to determine whether flow laminar or turbulent

### Laminar & Turbulent flow



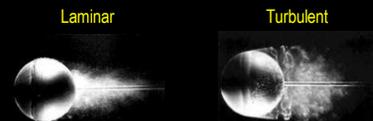
### Reynolds number

$$Re = \frac{\rho q d}{\mu}$$

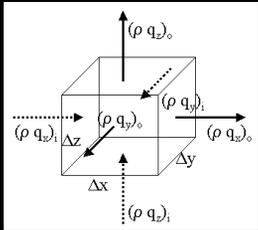
- Where
  - $\rho$  is density
  - $q$  is discharge velocity
  - $d$  diameter of discharge passageway
  - $\mu$  is the viscosity

### Characteristic Numbers

- If  $Re$  smaller than  $\sim 2200$  the fluid flow is laminar, if higher then the flow is turbulent.



## Control Volume



## Vatnafræði AQUIFERS

### Types

## What are Aquifers?

- Groundwater-bearing formations sufficiently permeable to transmit and yield water in usable quantities are called aquifers.
  - Most common are unconsolidated sands and gravels.
- Two types of aquifers: (1) unconfined and (2) confined.

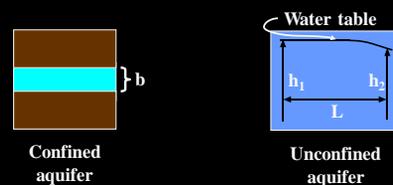
## Unconfined aquifers

- Are under-ground lakes in porous materials
- Top of the unconfined aquifer is the **water table**
  - Also called phreatic surface, after the Greek word phrear, "well"
- That is the plane where groundwater pressures are equal to atmospheric pressure.

## Confined Aquifer

- A layer of water-bearing material that is sandwiched between two layers of much less pervious material
- The pressure condition is characterized by the *potentiometric surface*, which is the surface obtained by connecting equilibrium water levels in tubes, or piezometers, penetrating the confined aquifer

## Aquifer types



### Potentiometric (Piezometric) Surface

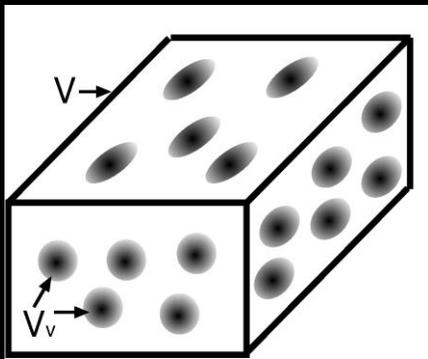
- If it is above the upper confining layer, the static water level in a well will be above the aquifer.
- Such a well is called an *artesian* well, named after wells bored in Artois (N-France) in the eighteenth century.

### Gropa / Porosity

- Hlutfall holrýmis á móti heildarrúmmáli  
– The percentage that is void of material,

$$n = \frac{V_v}{V} \times 100$$

- Þar sem  $V_v$  er holrýmið,  $V$  heildarrúmmálið og  $n$  gropan  
– Where  $V_v$  is the void space,  $V$  the total volume, and  $n$  porosity



### Calculating $n$

- $V_v = V - V_s$   
– where  $v$  – void,  $s$  – solid
- $V_s = m_s / \rho_s$   
– where  $\rho$  – density,  $m_s$  – dry weight

### Porosity of materials

- Well rounded and sorted sediments have porosity of 26% – 48%
- If well sorted, then porosity depends only on the packing (not grain size)

### Packing



Cubic packing  
of spheres

$$n = 47.65\%$$



Rhombohedral  
packing of spheres

$$n = 25.95\%$$

### Porosity – mixed grain sizes

- If a mixture of grain sizes, then porosity will be lower
  - Small grains fill in the void space
- Porosity of sedimentary rocks 1 – 30%
- Plutonic rock ~2% (+2 – 5% if fractured)
- Weathering can increase to 30 – 60%

### Effective porosity, $n_e$

- Measured by drying the sample at 105°C
- Then submerged in known volume of water
- What goes into sample is a measure of the effective void space

### Conductivity

- If the voids are poorly interconnected, the rock cannot convey water from one void to another.

### Hydraulic conductivity

- Also referred to as the  
**Coefficient of permeability**
- Intrinsic permeability,  $K_i = C d^2$ 
  - C – shape factor
  - d – grain size (diameter)

### Estimating water content

- **Specific yield,  $S_y$** 
  - Ratio of volume of water drained by gravity to total volume
- **Specific retention,  $S_r$** 
  - Ratio of volume of water a rock can retain to total volume
- If two rocks have **equal porosity**, but **different grain size**, **more water** will be retained in the **fine grained** rock, why?
  - Surface tension

### Specific yield and retention

- The sum of those is,

$$n = S_y + S_r$$

→ porosity

## Aquifer transmissivity

$$T = bK$$

- $T$  – transmissivity [ $L^2 T^{-1}$ ]
- $b$  – thickness of saturated layer [L]
- $K$  – hydraulic conductivity [ $L T^{-1}$ ]



## Water withdrawal

- Common to withdraw groundwater but too fast or extensive withdrawal (mining) can lead to:
  - Cone of depression
  - Wells can go “dry”
  - Lower water table
    - Serious consequences for vegetation
  - Increased concentration of chemicals in water
  - Saltwater intrusion
  - Subsidence

## Saltwater intrusion

- Saltwater flowing inland into aquifers
- Sea water has higher density than freshwater - and thus the pressure under saltwater is higher than with freshwater
- The withdrawal of freshwater and this pressure difference triggers a flow from the saltwater column to the freshwater column

## Subsidence

- Drained soil compacts
- Water table falls, peat exposed to oxygen
- Decomposes
- Subsidence

