

The Antarctic contribution to Holocene global sea level rise

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The Holocene glacial and climatic development in Antarctica differed considerably from that in the Northern Hemisphere. Initial deglaciation of inner shelf and adjacent land areas in Antarctica dates back to between 10–8 Kya, when most Northern Hemisphere ice sheets had already disappeared or diminished considerably. The continued deglaciation of currently ice-free land in Antarctica occurred gradually between ca. 8–5 Kya. A large southern portion of the marine-based Ross Ice Sheet disintegrated during this late deglaciation phase. Some currently ice-free areas were deglaciated as late as 3 Kya. Between 8–5 Kya, global glacio-eustatically driven sea level rose by 10–17 m, with 4–8 m of this increase occurring after 7 Kya. Since the Northern Hemisphere ice sheets had practically disappeared by 8–7 Kya, we suggest that Antarctic deglaciation caused a considerable part of the global sea level rise between 8–7 Kya, and most of it between 7–5 Kya. The global mid-Holocene sea level high stand, broadly dated to between 8–4 Kya, and the Littorina-Tapes transgressions in Scandinavia and simultaneous transgressions recorded from sites e.g. in Svalbard and Greenland, dated to 7–5 Kya, probably reflect input of meltwater from the Antarctic deglaciation.

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Introduction

The modern Antarctic Ice Sheet (Fig. 1) covers an area of ca. $13.6 \times 10^6 \text{ km}^2$, and consists of roughly $29 \times 10^6 \text{ km}^3$ of grounded ice, with a total potential effect on sea level rise of about 66 m (Drewry et al. 1982; Denton et al. 1991). The role of Antarctic glacial ice as a meltwater source for global sea level rise since the Last Glacial Maximum (LGM) is controversial (Andrews 1992). Estimates range from as little as 0.5–2.5 m (Colhoun et al. 1992) to 37 m (Nakada & Lambeck 1988), depending on whether calculations are based on minimum or maximum reconstruction of LGM-ice volumes in Antarctica. A number of appraisals put the Antarctic contribution to global sea level rise since LGM on the order of 12–26 m (Hughes et al. 1981; Huybrechts 1990a, b; Tushingham & Peltier 1991; Peltier 1994; Clark et al. 1996). This is out of the total estimated global sea level rise of 120 ± 20 m since LGM (e.g. Chappell & Shackleton 1986; Fairbanks 1989; Peltier 1994).

The timing of LGM in Antarctica has not been precisely determined. It has been radiocarbon

dated to 20–17 Kya (thousands of years before present) in the western Ross Sea area (Stuiver et al. 1981; Anderson et al. 1992; Kellogg et al. 1996; Licht et al. 1996), to 14 Kya in Taylor Valley, southern Victoria Land (Hall 1997), and to ≥ 17 and ≥ 14 Kya off East Antarctica and the Antarctic Peninsula, respectively (Banfield & Anderson 1995; Harris et al. in press). Most authors assume that LGM roughly coincides with the lowest global sea level stand at 20–18 Kya.

In order to make a fair estimate of the Antarctic contribution to global sea level rise since LGM, the overall ice volume at LGM needs to be established. However, existing appraisals are controversial. One category of estimates suggests that marginal domes of the Antarctic Ice Sheet were 500–1000 m thicker than at present and that ice extended to the shelf break around most of the continent (Clark & Lingle 1979; Denton 1979; Hughes et al. 1981; Zhang 1992). Hughes et al. (1981) calculated the volume of grounded ice in Antarctica at LGM to have been $37 \times 10^6 \text{ km}^3$, compared to roughly $29 \times 10^6 \text{ km}^3$ today, and suggested that the difference of $8 \times 10^6 \text{ km}^3$ equalled about 24 m of global sea

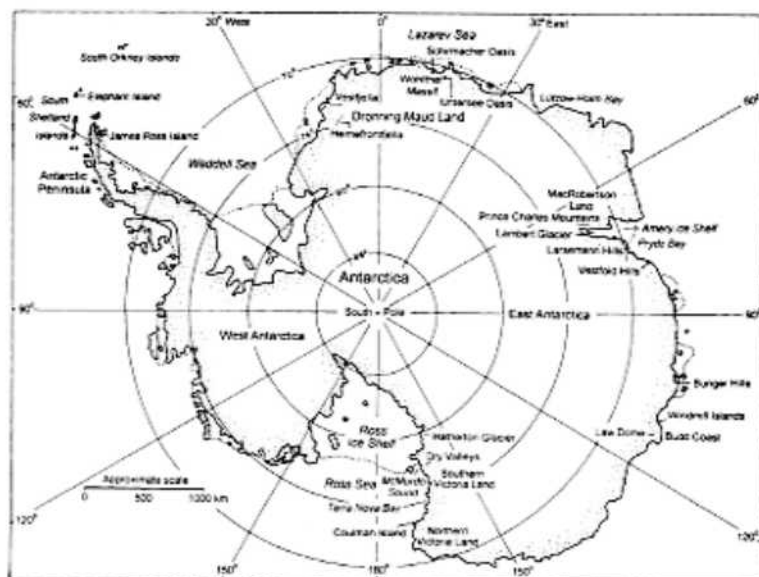


Fig. 1. Map of Antarctica.

level lowering at LGM. Their reconstruction was based on the assumptions that the East Antarctic Ice Sheet's peripheral domes thickened considerably and expanded towards the shelf break, and that the West Antarctic Ice Sheet was considerably expanded relative to its present configuration. Some other estimates advocate a considerably smaller ice extent at LGM, maybe only insignificantly larger than today's in some parts of Antarctica (Mayewski 1975; Colhoun & Adamson 1992; Colhoun et al. 1992; Goodwin 1993; Hayashi & Yoshida 1994). A recent review of Antarctic glacial history since the LGM (Ingólfsson, Hjort, Berkman et al. 1998) suggests that while ice extended well offshore around most of Antarctica at LGM, it did not extend all the way to the shelf break off East Antarctica and in the western Ross Sea, and that some East Antarctic coastal oases were probably ice-free. One interpretation of Antarctic ice-core data suggests that during LGM accumulation rates in interior Antarctica were 50–60% lower than mean Holocene rates (Raisbeck et al. 1987; Jouzel et al. 1989), and that LGM ice thickness in interior East Antarctica was less than during the Holocene (Raynaud & Lebel 1979; Martinerie et al. 1994).

Lacking good estimates of Antarctic ice volumes at LGM, an alternative way to examine the Antarctic contribution to post-LGM sea level rise is to look at the chronology of deglaciation and compare that with the Northern Hemisphere devel-

opment. If the deglaciation history of Antarctica differs substantially from that in the north, the Antarctic contribution might show up as a water-volume input causing global sea level rise at a time when contribution from Northern Hemisphere ice sheets was unlikely. Thus a minimum estimate of the Antarctic contribution should be possible.

The Northern Hemisphere deglaciation

The Laurentide and Cordilleran ice sheets

The deglaciation of the North American ice sheets mainly took place between 16–8 Kya (Dyke & Prest 1987a; Fulton 1989; Lundqvist & Saarnisto 1995). The Cordilleran Ice Sheet retreated rapidly after 15 Kya, and disappeared between 10–8 Kya (Dawson 1992; Lundqvist & Saarnisto 1995). According to Dyke & Prest (1987a, b), the Laurentide Ice Sheet began retreating between 17–14 Kya. Between 13–11 Kya ice retreat and thinning accelerated. Huge proglacial lakes, the most important of which was Lake Agassiz, developed, and large volumes of meltwater were discharged into the Gulf of Mexico (Teller 1990). Between 11–10 Kya ice retreat caused the drainage of Lake Agassiz to change, so that its meltwater was discharged into the western North Atlantic. By ca. 8 Kya the Laurentide Ice Sheet disintegrated over Hudson Bay and the last documented major

meltwater input through the Hudson Strait into the North Atlantic was a huge drainage event from glacial lakes Ojibway and Agassiz in connection with the Hudson dome collapse (Dyke & Prest 1987a; Andrews 1989, 1994). After 8 Kya much of the ice retreat was associated with ice stagnation and by 7.2–7 Kya only small remnants of the Laurentide Ice Sheet still existed in Labrador and on Baffin and Melville islands (Dyke & Prest 1987a, b, c; Andrews, Mclean et al. 1995). Likewise, only small remnants of the Inuit ice sheet of north-eastern Arctic Canada remained by 8–7 Kya. A comparison between the deglaciation history and the regional $\delta^{18}\text{O}$ isotope record from the south-eastern Baffin Island shelf suggests a rapidly decreasing meltwater input to Hudson Bay and Labrador Sea after 8 Kya (Andrews, Erlenkeuser et al. 1991).

The ice sheets of Greenland and Iceland

The Late Weichselian–early Holocene glacial history of Greenland was recently reviewed by Funder & Hansen (1996). They suggested that the initial retreat of the Greenland Ice Sheet after LGM was caused by rising global sea level, destabilizing marine portions of the ice sheet. The period between 10–9.5 Kya was characterized by unstable ice margins, with glacial re-advances in many parts of the country (Funder et al. 1998). However, after 9.5 Kya the ice margins retreated rapidly, and by 8–7 Kya the present distribution of ice cover was attained around Greenland (Funder 1989). The Icelandic Ice Sheet retreated to positions inside the present coast already by 13–12 Kya, but experienced considerable ice re-advances between 11–10 Kya (Ingólfsson & Norddahl 1994). After 10 Kya the Icelandic Ice Sheet retreated and thinned very rapidly, and by 8 Kya BP the Icelandic glaciers were of a similar size or smaller than today (Ingólfsson, Björck et al. 1997).

The Scandinavian–British Ice Sheet

In Scandinavia and the British Isles ice thinning and retreat was under way by 16 Kya (Dawson 1992), and the deglaciation of the British Isles was complete by ca. 10 Kya (Walker et al. 1994). By 10 Kya coastal Norway, southern Sweden and south-easternmost Finland were deglaciated (Birks et al. 1994; Lundqvist & Saarnisto 1995). After 10 Kya the Scandinavian inland ice retreated and thinned very rapidly. By 9 Kya remnants of the ice

sheet lingered in the Norwegian–Swedish mountain range and on the lowlands of central northern Sweden, but by 8.5 Kya only minor ice fields remained in the high mountains of northern Sweden (Lundqvist & Saarnisto 1995).

The Barents Sea and Kara Sea ice sheets

There is good evidence for an ice sheet grounded over much of the Barents Sea shelf during the Late Weichselian (Elverhøi & Solheim 1983; Polyak et al. 1995). The Barents Sea Ice Sheet was centred over the northern Barents Sea and eastern Svalbard (Salvigsen 1981; Forman 1990), but probably expanded across Spitsbergen and terminated beyond the present coastline of Svalbard (Forman 1989; Mangerud et al. 1992; Landvik, Bondevik et al. 1998). Recent studies suggest relatively limited ice volumes over Novaya Zemlya and the Kara Sea during the Late Weichselian (Forman, Lubinski et al. 1995; Astakhov 1998; Svendsen et al. 1999; Forman, Ingólfsson et al. 1999). Deglaciation from the shelf areas was under way by 14–13 Kya, and by 10–9.5 Kya these marine-based ice sheets had completely disintegrated. Glaciers on Svalbard were at or behind their present margins before 9 Kya (Salvigsen 1981; Landvik, Bondevik et al. 1998).

Small ice caps and mountain glaciers

The combined volume of ice in glaciers and small ice caps in the mountain ranges of Europe, North America, the Himalayas, South American Andes, as well as in New Zealand, Tasmania and in the high mountains in Africa and New Guinea during the last glacial maximum has been estimated to equal a global sea level change of about 5.5 m (Hughes et al. 1981). Since most of these had melted away already by 10 Kya (Denton & Hughes 1981; Clapperton 1990; Dawson 1992; Rutter 1995) or shortly thereafter, they did not significantly contribute to global sea level rise in Holocene times.

The Antarctic deglaciation

There have been different views as to whether the reduction of the Antarctic Ice Sheet since the LGM preceded, was in phase with, or lagged behind the Northern Hemisphere deglaciation (Nakada & Lambeck 1988; Colhoun et al. 1992; Hjort, Ingólfsson et al. 1997). The prevailing view has been that the East Antarctic Ice Sheet remained

relatively stable during the transition from LGM to the Holocene, whereas the West Antarctic Ice Sheet experienced some major changes (Stuiver et al. 1981), and further, that fluctuations of the West Antarctic Ice Sheet were broadly synchronous with ice sheets in the Northern Hemisphere (Clapperton & Sugden 1988; Broecker & Denton 1990; Clapperton 1990; Dawson 1992). These views have been increasingly challenged by results coming from different regions around Antarctica (e.g. Barsch & Mäusbacher 1986; Domack et al. 1991; Ingólfsson, Hjort, Björck et al. 1992; Goodwin 1993; Verkulich & Hiller 1994; Björck, Håkansson et al. 1996; Björck, Hjort et al. 1996; Licht et al. 1996; Shevenell et al. 1996; Colhoun 1997; Hjort, Ingólfsson et al. 1997; Hjort, Björck et al. 1998; Zwartz et al. 1998). Recently Ingólfsson, Hjort, Berkman et al. (1998) examined the circumantarctic glacial history evidence on land, related to developments after the LGM. The purpose was to check if a general Antarctic glacial history pattern could be identified, and to highlight the question of the relative timing of deglaciation in Antarctica versus that of the large Northern Hemisphere ice sheets. They recognized a broad pattern in the glacial histories from the coastal sites around Antarctica (Fig. 2):

a) Ice extended offshore around most of Antarctica at LGM, though some oases may have remained ice free. But the LGM ice sheet did not reach the shelf break off East Antarctica or in the western Ross Sea or along the southernmost parts of the Antarctic Peninsula. Ice retreat from the LGM positions was under way by 17–14 Kya, and by 11–10 Kya initial deglaciation of some inner shelf areas as well as a few outer coastal land areas had occurred. The ice retreat from LGM positions around Antarctica was probably eustatically controlled. Melting Northern Hemisphere glaciers caused global sea level to rise and the Antarctic grounding lines and ice fronts gradually retreated from the outer shelf towards the continent.

b) Deglaciation of some shallow inner shelf areas and of most presently ice-free land areas, with the exception of some East Antarctic oases, occurred between ca. 10–5 Kya. After the receding glaciers had reached the edges of most of the presently ice-free land areas or shallow inner shelf areas at 10–8 Kya, the rate of deglaciation slowed – and sometimes reversed into glacial re-advances. It was not until around or shortly after 5 Kya that most Antarctic glaciers had retreated to or behind

their present positions. Some areas seem to have become ice-free as late as ca. 3 Kya, and the grounding line of the Ross Ice Shelf continued retreating until ca. 2.5 Kya.

c) This gradual deglaciation between 10–5 Kya was most likely controlled by a globally seen delayed interglacial warming in Antarctica. The warming peaked late, in a hypsithermal event visible in the records between 4.7–2 Kya. The best documented records put it between 4–3 Kya, in the Antarctic Peninsula region, and between 3.6–2.6 Kya in coastal Victoria Land. Many records

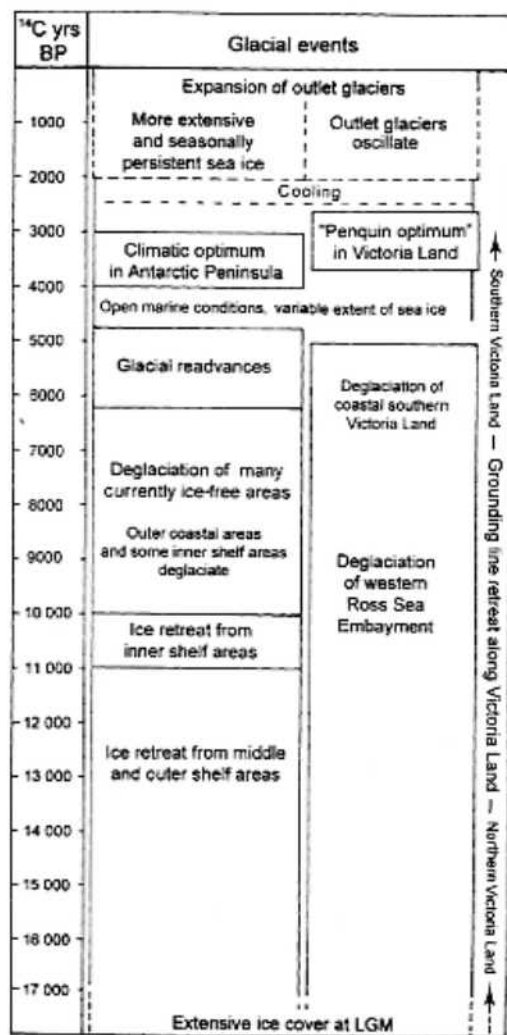
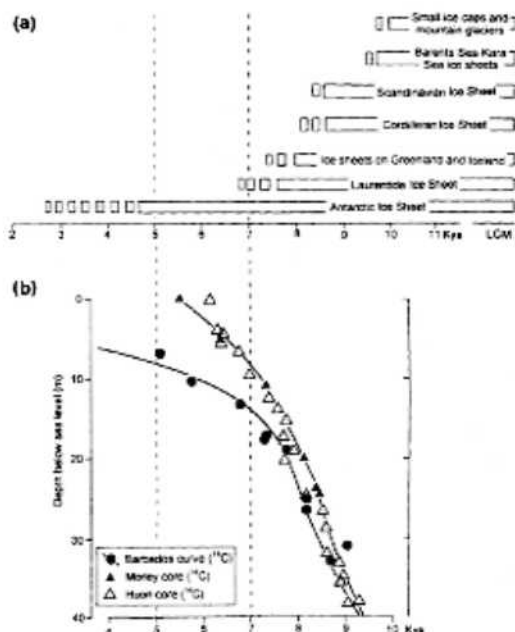


Fig. 2. A broad synthesis of the deglaciation of Antarctica and associated environmental developments since the LGM (slightly modified from Ingólfsson, Hjort, Berkman et al. 1998).

Fig. 3. (a) Timing of deglaciation and meltwater input to the oceans from the Northern Hemisphere and Antarctic ice sheets. Whole columns indicate substantial input; broken columns indicate minor input. Northern Hemisphere input diminishes rapidly after 10 Kya, and by 8 Kya only the Laurentide Ice Sheet continues to contribute significant meltwater. After ca. 7 Kya the Antarctic Ice Sheet is the only plausible major meltwater source. (b) Sea level rise after 9 Kya, according to the Barbados sea level curve of Fairbanks (1989) and the sea level data from Australian Morley Island core (Eisenhauer et al. 1993) and the Papua New Guinea Huon Peninsula core (Chappell & Polach 1991). Although large (ca. 10 m) differences of apparent sea level depth at 6 Kya can be observed between the Australian and New Guinea data and the Barbados curve, the combined data indicate considerable sea level rise between ca. 8–5 Kya. The differences in apparent sea level depth have been explained as resulting from water redistribution due to changes in the Earth's geoid after 6 Kya (Eisenhauer et al. 1993). Period between 7–5 Kya, when the Littorina-Tapes transgressions are recorded in Scandinavia and where the Antarctic meltwater input shows strongly, is marked with broken lines between (a) and (b).



indicate less sea ice and higher beach energies in the mid-Holocene, which probably coincide with this warm event, than at present.

d) This hypsithermal period was followed by a general cooling after 2.5 Kya, and expanding outlet glaciers in many areas around Antarctica.

Discussion and conclusions

We see distinct differences between the deglaciation histories of the large Northern Hemisphere ice sheets and Antarctica (Fig. 3). Between 10–8 Kya the meltwater input from the Northern Hemisphere ice sheets diminished rapidly, as they disintegrated and melted away. There were minimal or no meltwater inputs from the Cordilleran Ice Sheet and the Barents Sea-Kara Sea ice sheets after 10–9.5 Kya. The Scandinavian Ice Sheet continued contributing meltwater until ca. 8.5 Kya. Meltwater input from the Greenland Ice Sheet was minimal after 8 Kya. The Laurentide Ice Sheet contributed large volumes of meltwater until ca. 8 Kya, but its input dwindled rapidly between 8–7 Kya and was minimal after 7 Kya. The global glacio-eustatic sea level rise, however, continued until some time between 6–5 Kya. As the terrestrial field evidence from Antarctica suggests that deglaciation there was distinctly delayed com-

pared to the Northern Hemisphere deglaciation, we suggest that this late Antarctic deglaciation caused part of the global sea level rise between 8–7 Kya and most of it between 7–5 Kya (Fig. 3).

Well-dated relative sea level records, e.g. from Scandinavia, Svalbard and Greenland, show marine transgressions in the broad interval 8–4 Kya (Hjort 1973, 1981; Forman 1990; Pirazzoli 1991; Donner 1995). Raised beaches of the Littorina-Tapes transgression of Scandinavia define a zone of high relative sea level, thought to have peaked at 7–6 Kya (Møller 1986; Donner 1995). In Svalbard, the mid-Holocene sea level record is characterized by a transgression, broadly dated between 6.5–5 Kya (Forman, Mann et al. 1987; Landvik, Mangerud et al. 1987; Forman 1990). Hjort (1973, 1981) described the Vega transgression in East Greenland, culminating shortly after 6 Kya, and Mid-Holocene transgressions or peaks in relative sea levels, broadly dated between 7–4 Kya, have been reported from many sites around the globe, both inside and outside the former glaciated areas (Pirazzoli 1991). A relative sea level fall of 1–2 m has been recorded at many sites since the mid-Holocene sea level high stand (Pirazzoli 1991). This relative sea level lowering has been attributed to the visco-elastic response of the Earth's mantle following increased water-load on the sea floor and the redistribution of the

